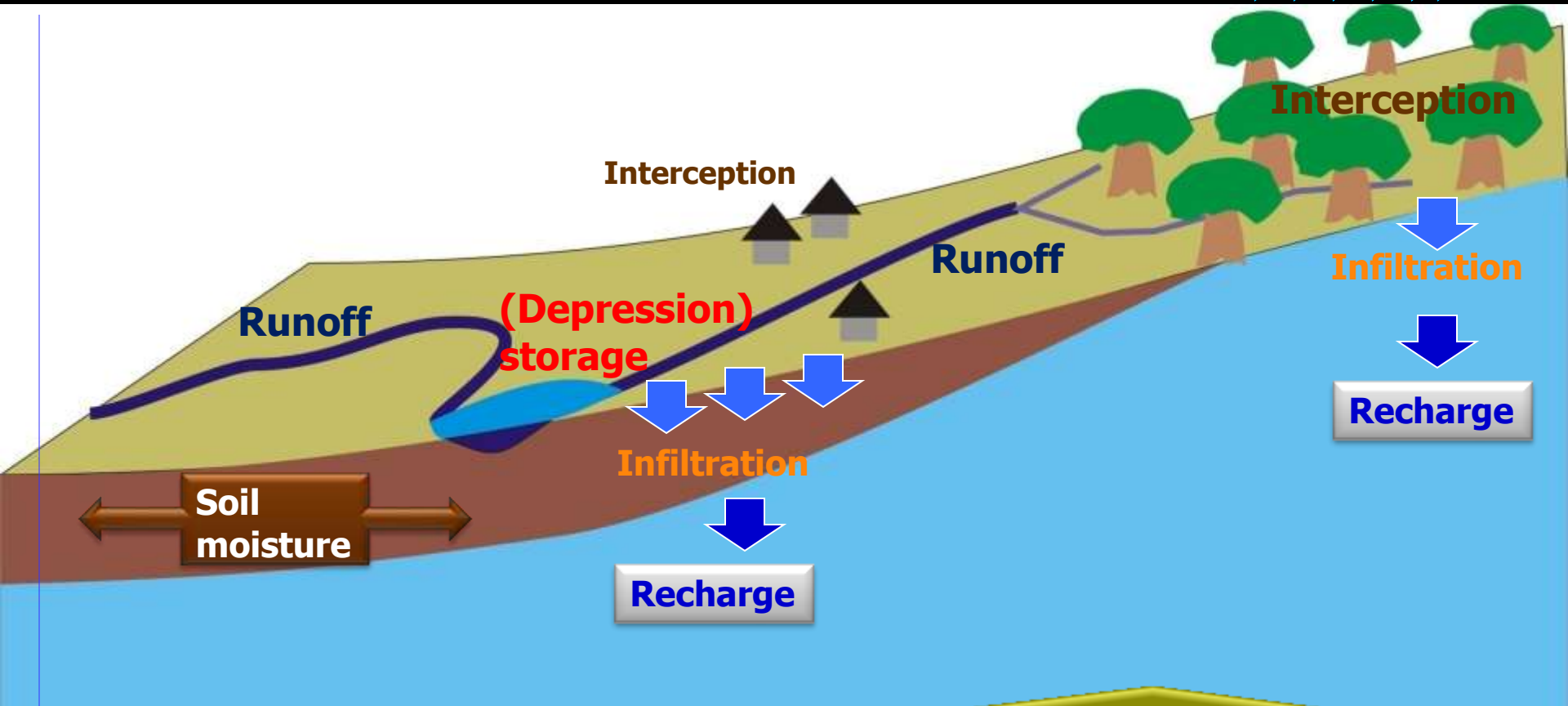


GROUNDWATER BALANCE

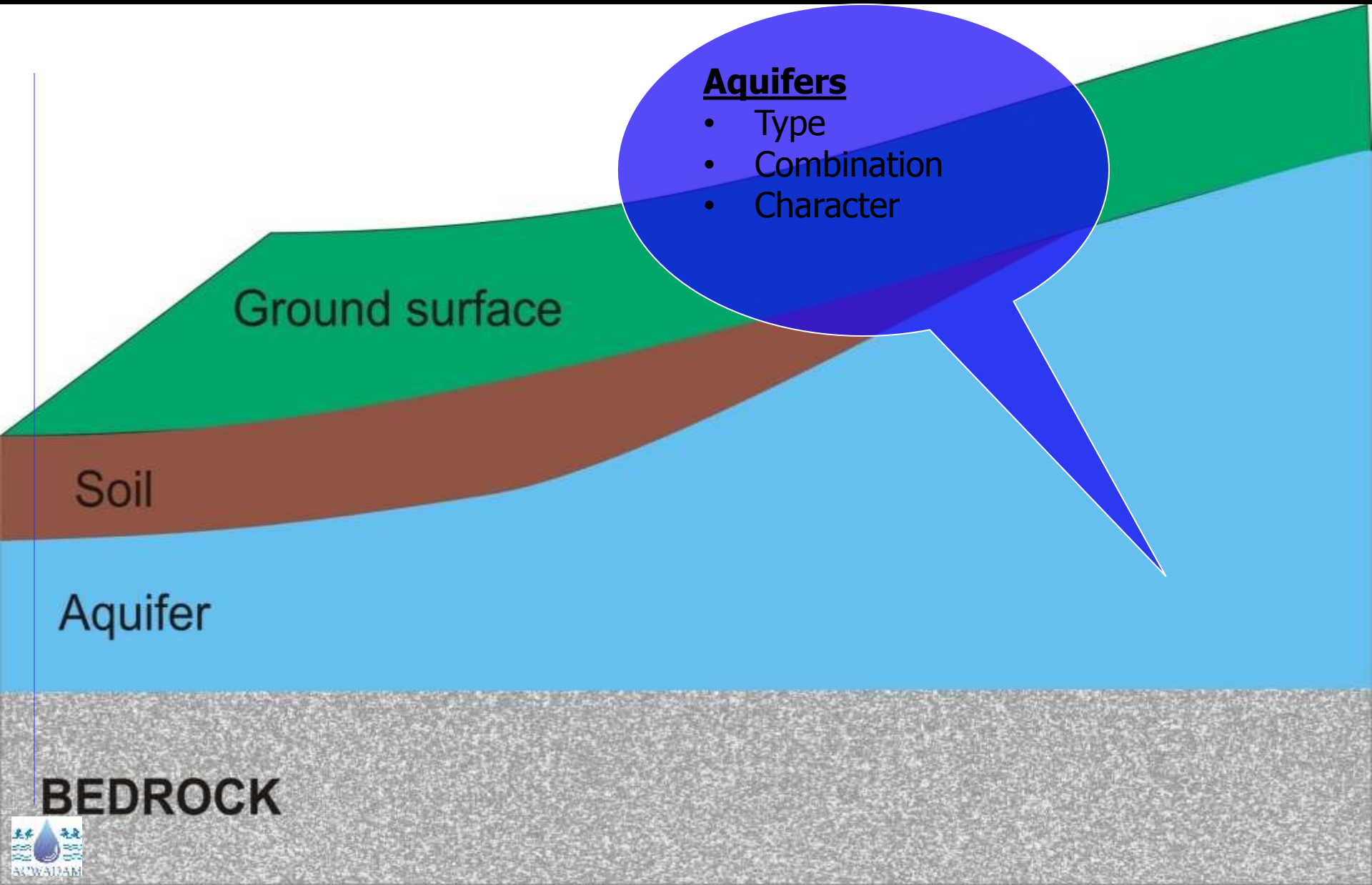


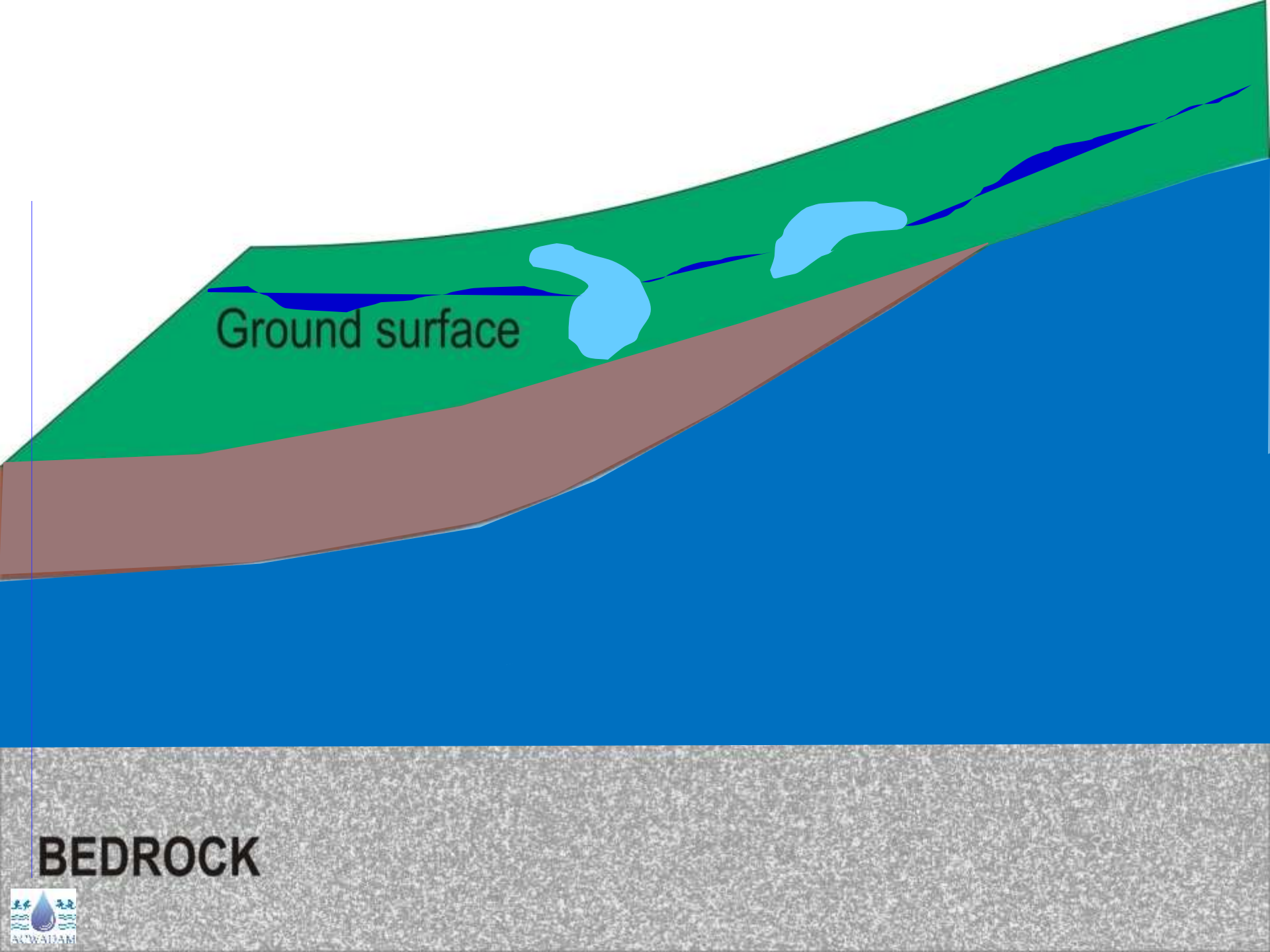
PRECIPITATION



$$\text{Precipitation} = \text{Interception} + \text{Evaporation} + \text{Transpiration} + \text{Runoff} + \text{Infiltration}$$

Groundwater balance...*aquifer as the unit*





Ground surface

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The water balance for an aquifer (groundwater balance) may vary, depending upon the nature of a groundwater system, say in the watershed

- Watershed with a deep aquifer
- Watershed has both shallow and deep aquifers
- Watershed has only shallow aquifer



Deep aquifer only

$$\text{Precipitation} = S_r + I + E_t$$

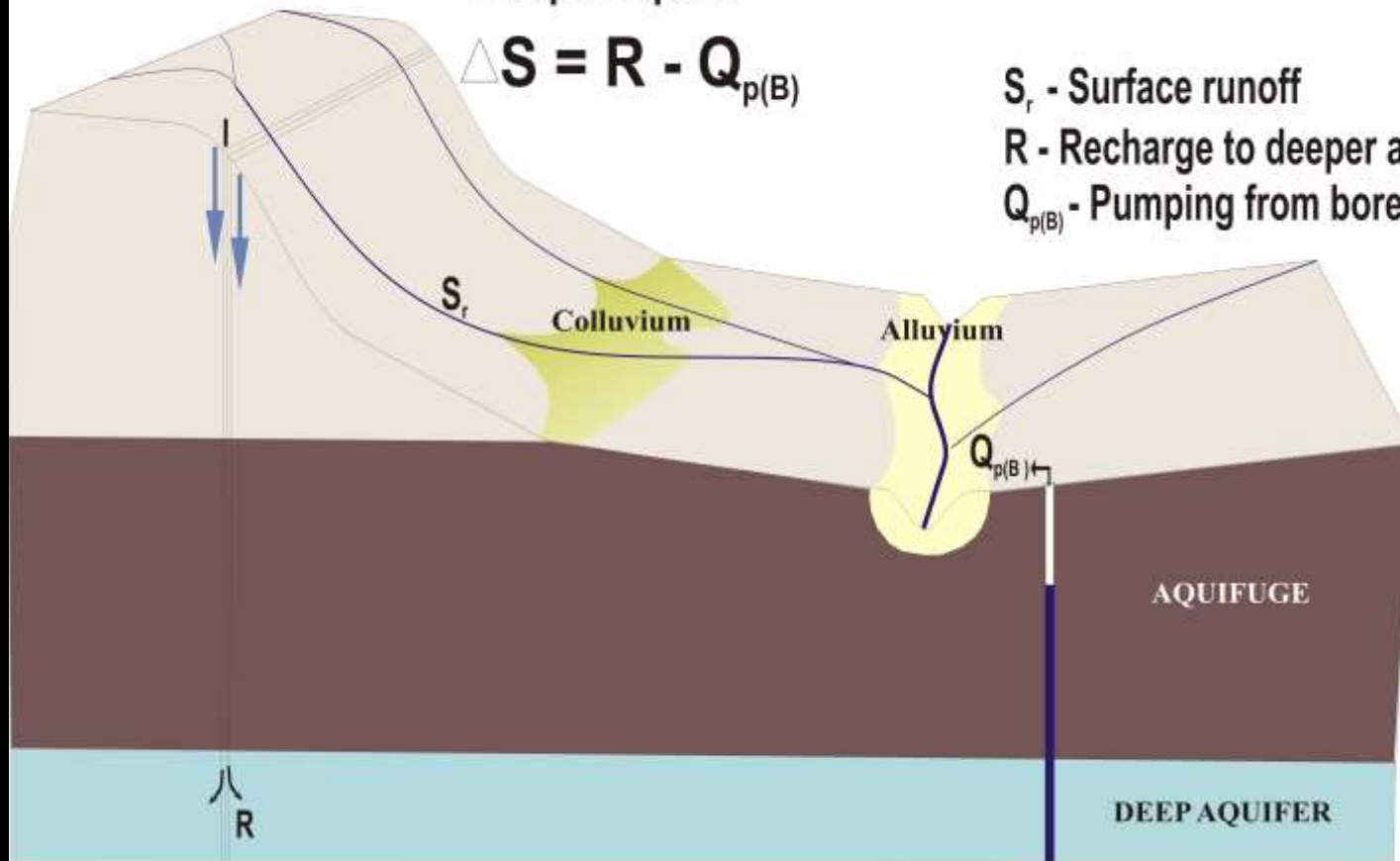
Deeper aquifer

$$\Delta S = R - Q_{p(B)}$$

S_r - Surface runoff

R - Recharge to deeper aquifer

$Q_{p(B)}$ - Pumping from bore wells in deeper aquifer



Shallow and deep aquifers

$$\text{Precipitation} = S_r + P + I + E_t$$

Shallow aquifer $\Delta S = I + Rf_i - Q_{p(D)} - B_f - L$

Deeper aquifer $\Delta S = L - Q_{p(B)}$

S_r - Surface runoff

E_t - Evapotranspiration

P - Percolation from plateau

I_p - Point source infiltration through deep fractures

I_d - Diffuse source infiltration along gentle slopes and plains

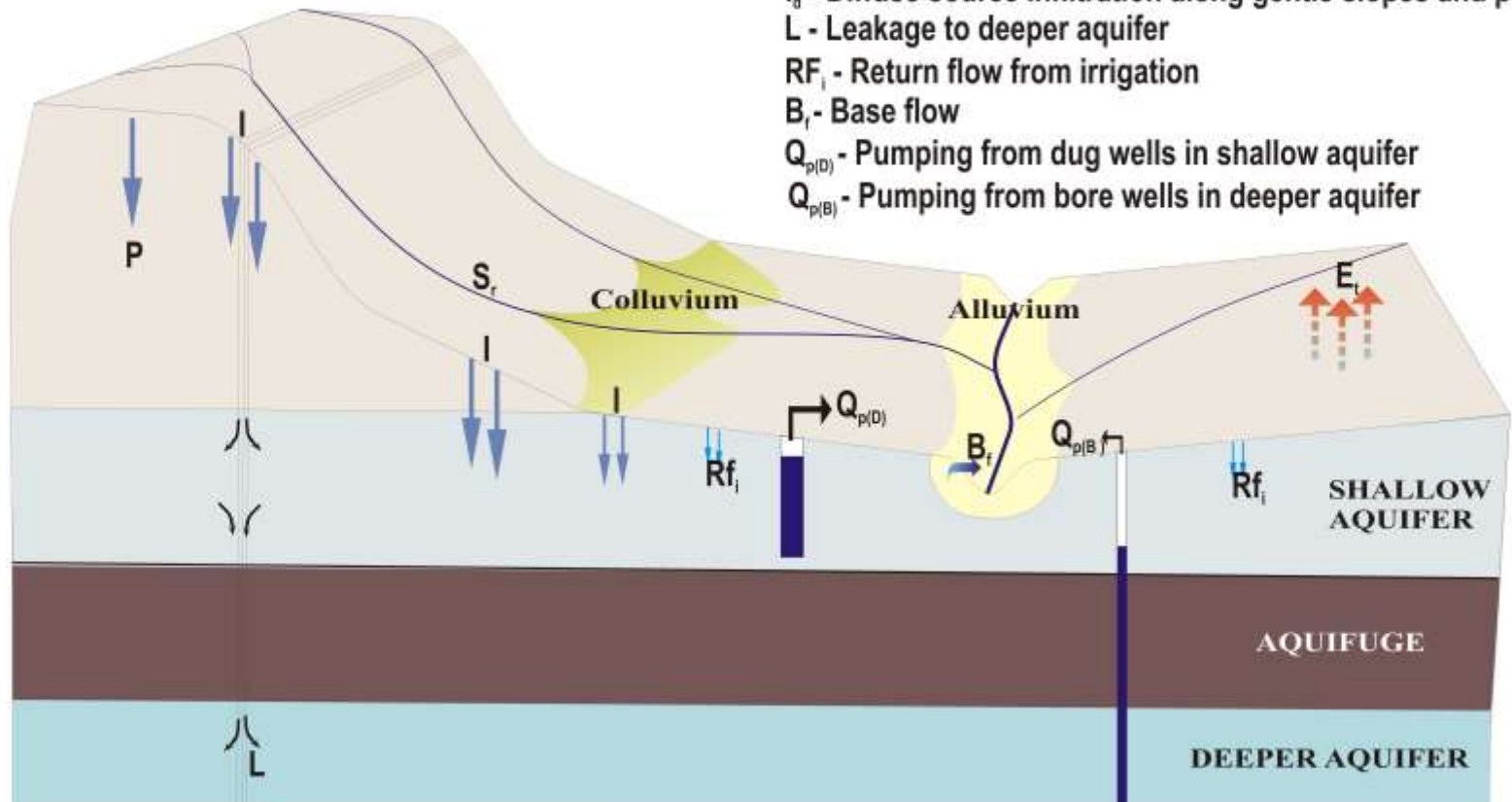
L - Leakage to deeper aquifer

RF_i - Return flow from irrigation

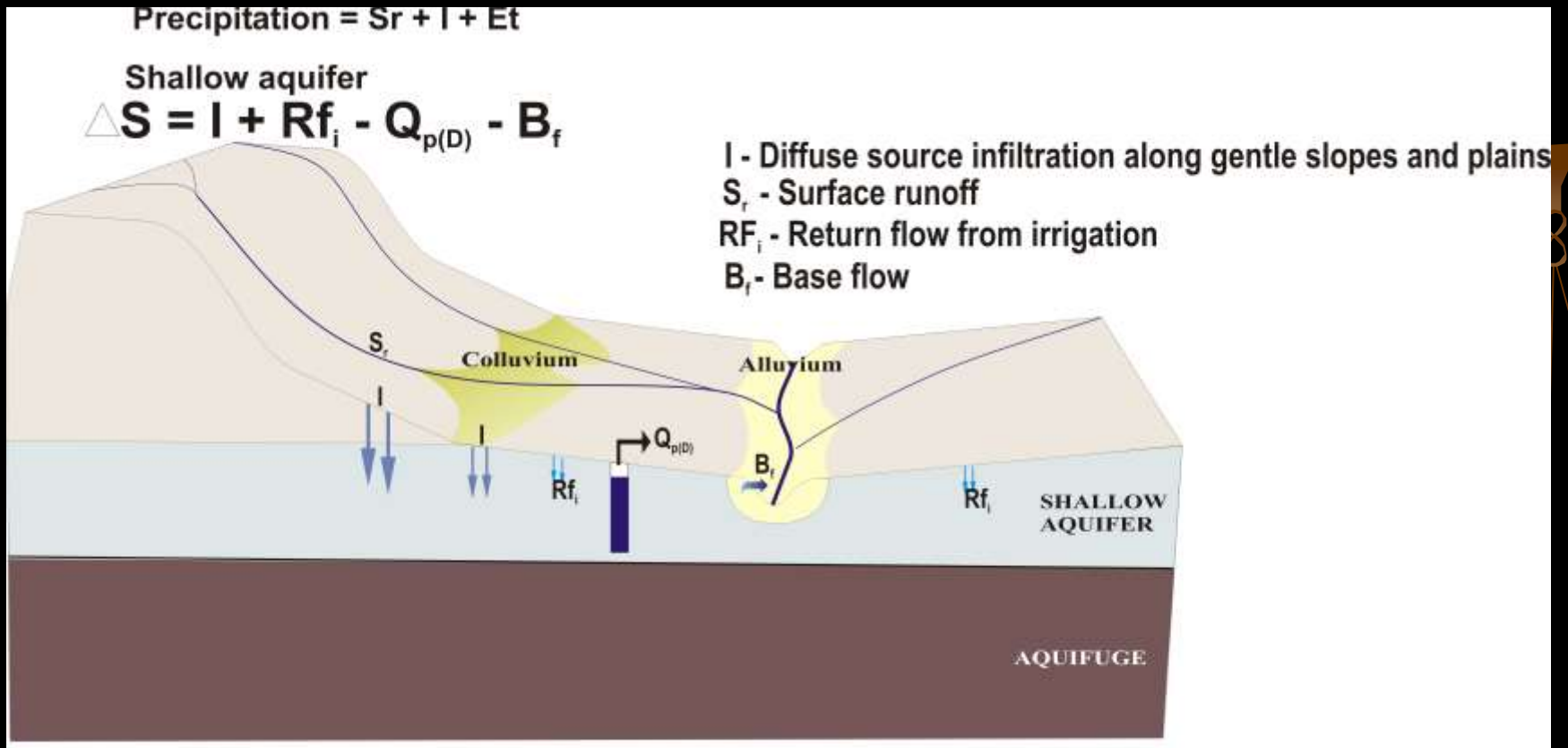
B_r - Base flow

$Q_{g(D)}$ - Pumping from dug wells in shallow aquifer

$Q_{p(B)}$ - Pumping from bore wells in deeper aquifer



Components of a shallow aquifer



An understanding of a groundwater system or a set of aquifers in a watershed requires:

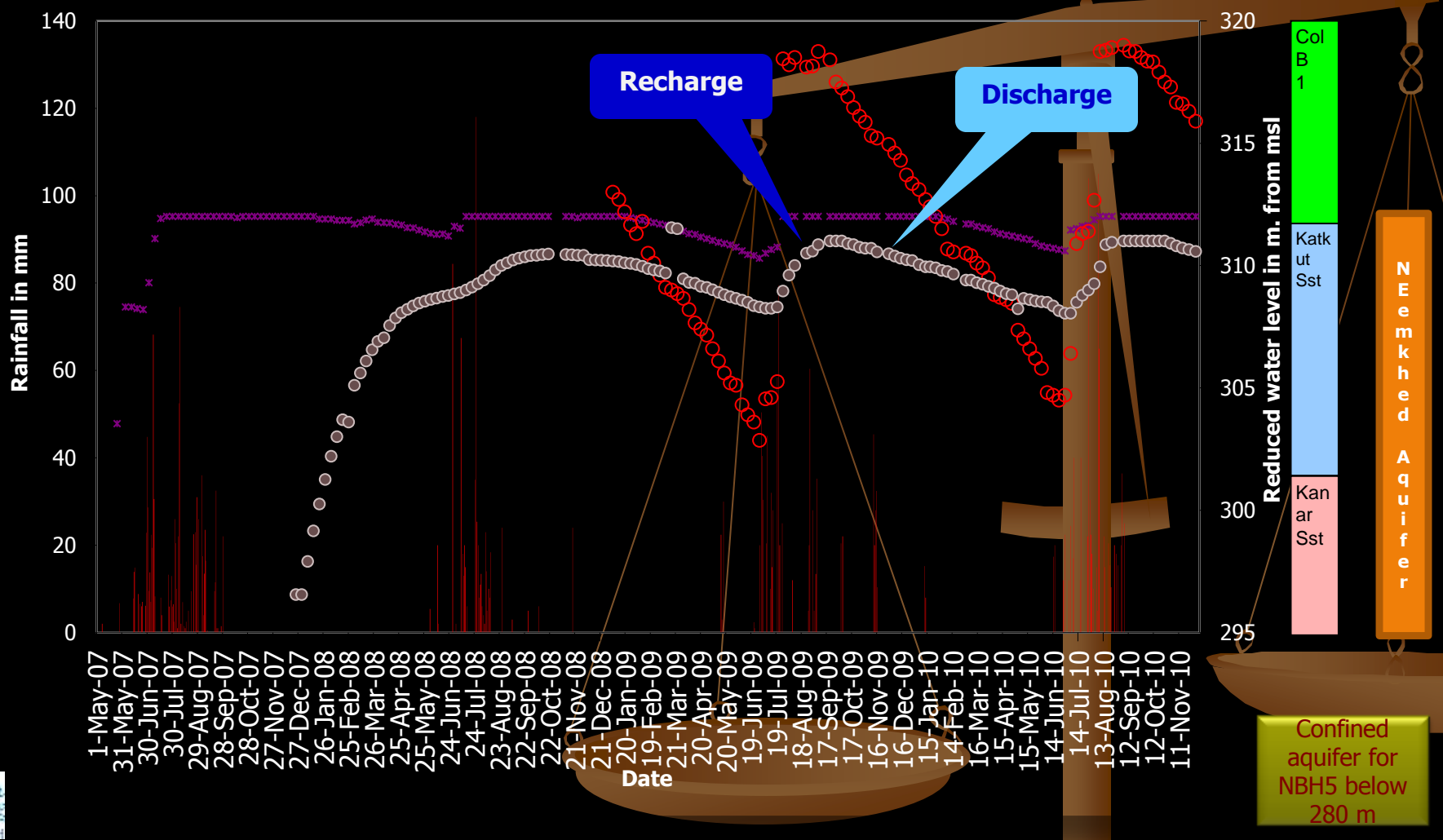
- Quantitative assessment of the INPUTS (Recharge) to and OUTPUTS (Discharge) from the aquifer.
- An idea about how and when these components work...

Rainfall

NBH6

Neemkheda dam

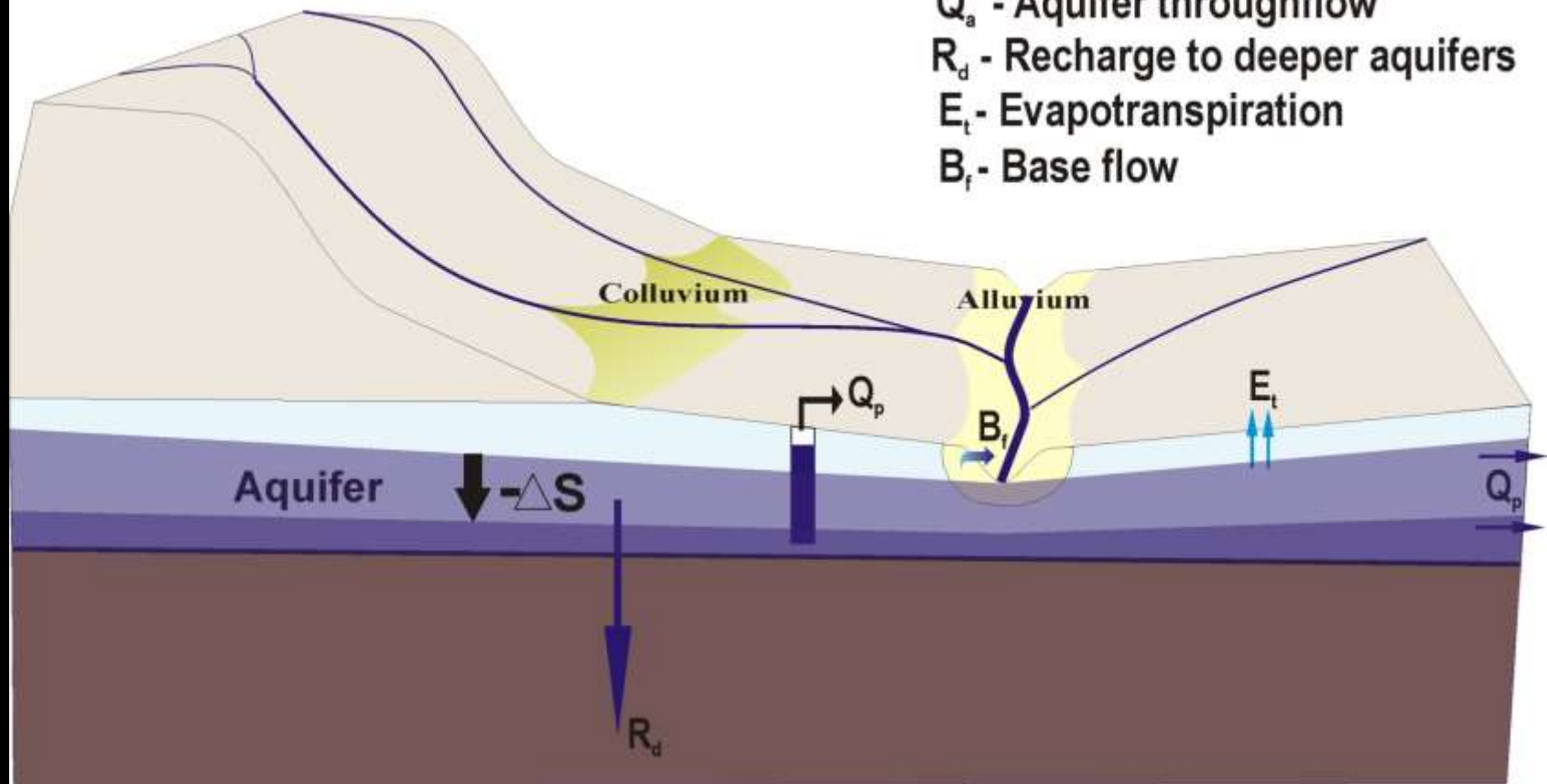
NBH5



Simplified groundwater balance equation

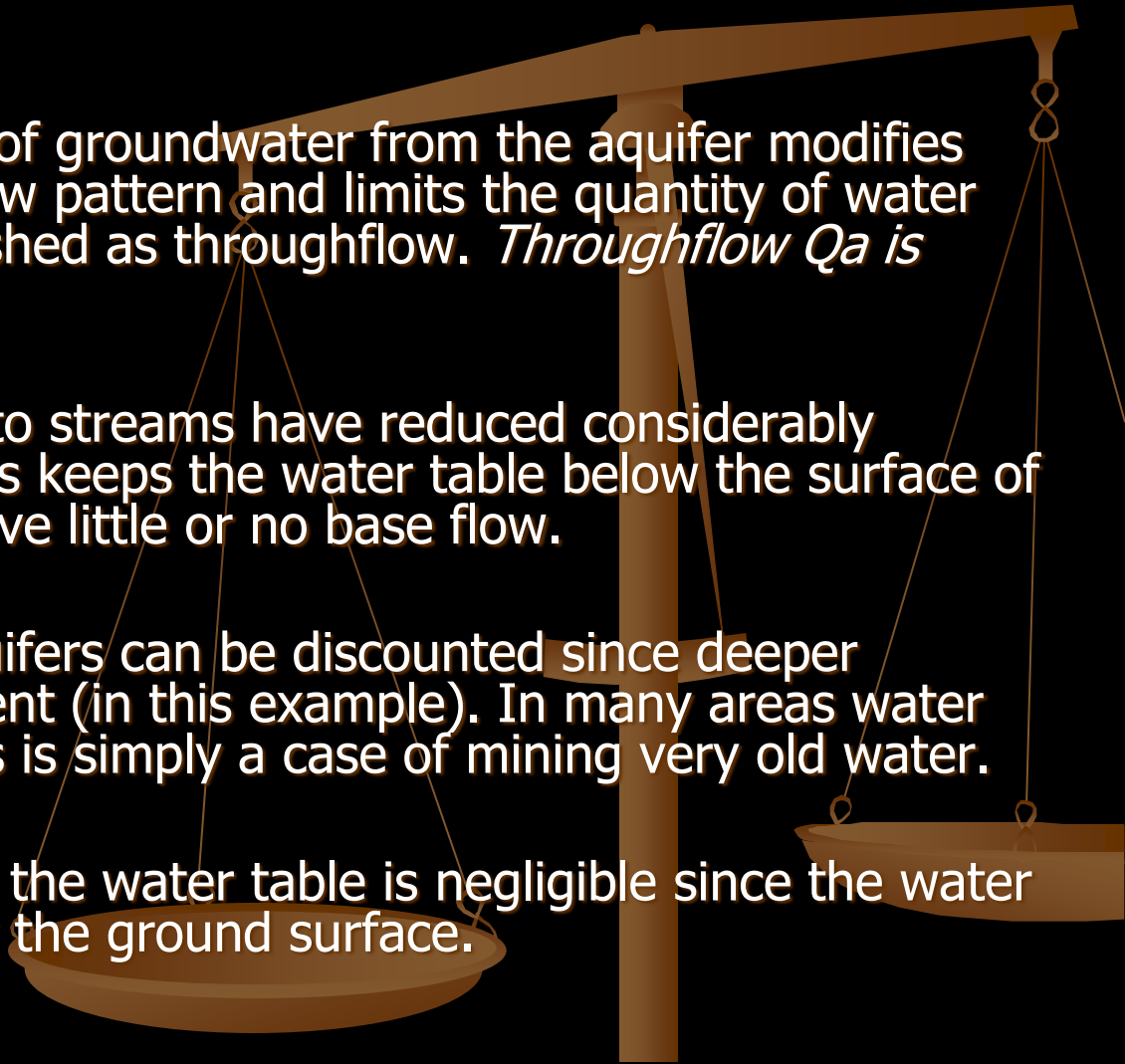
$$-\Delta S = Q_p + Q_a + R_d + E_t + B_f$$

Q_p - Groundwater pumping
 Q_a - Aquifer throughflow
 R_d - Recharge to deeper aquifers
 E_t - Evapotranspiration
 B_f - Base flow

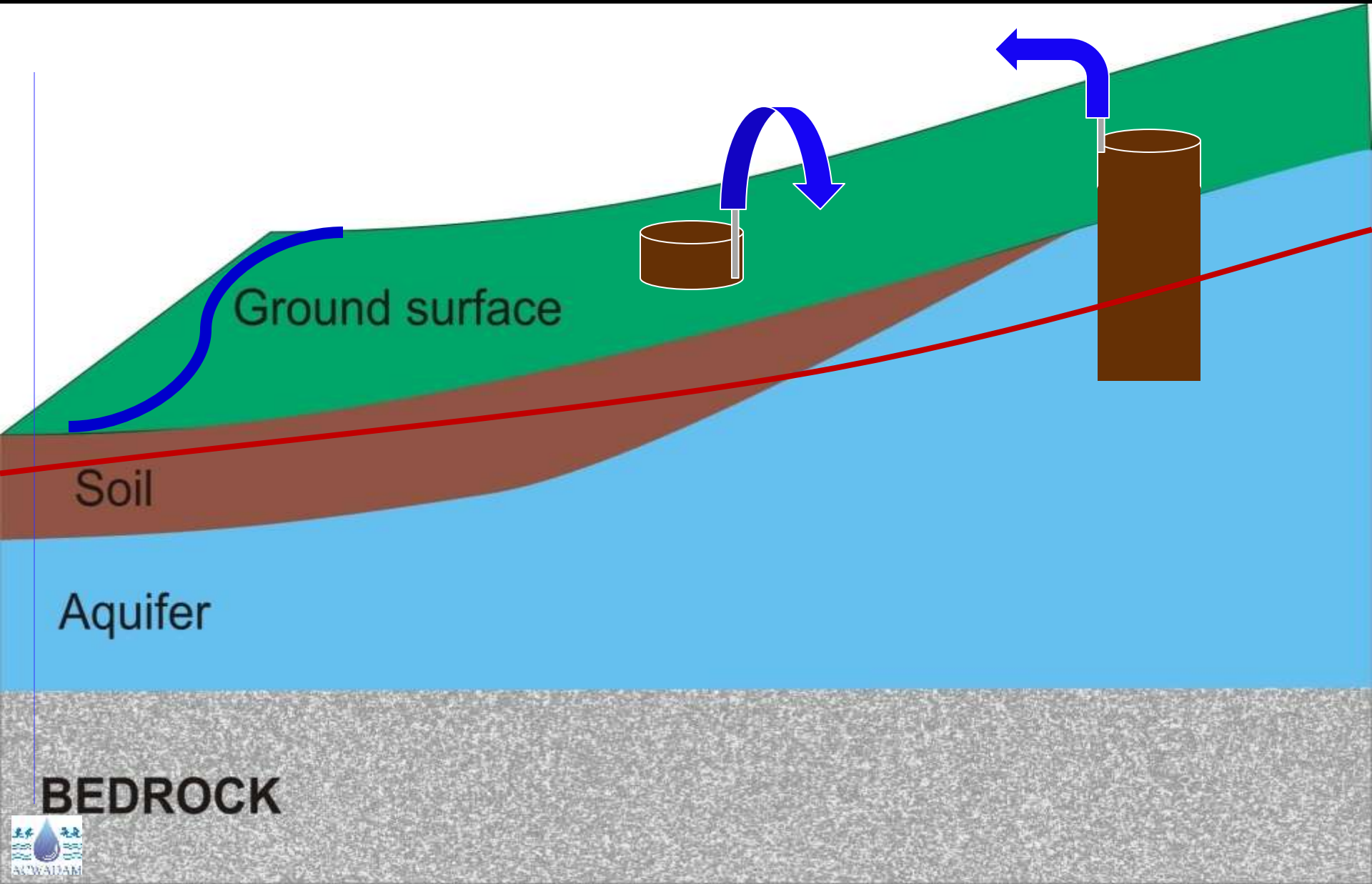


In many parts of India, where groundwater is pumped out in large quantities from shallow aquifers...and where water is limited in deeper aquifers

- Q_p - Groundwater pumping is the single largest component of groundwater discharge.
- Q_a - The intensive pumping of groundwater from the aquifer modifies the natural groundwater flow pattern and limits the quantity of water that flows out of the watershed as throughflow. *Throughflow Q_a is negligible or zero.*
- B_f - Baseflow contributions to streams have reduced considerably because pumping from wells keeps the water table below the surface of the streams that tend to have little or no base flow.
- R_d - Recharge to deeper aquifers can be discounted since deeper aquifers are practically absent (in this example). In many areas water drawn from deeper aquifers is simply a case of mining very old water.
- E_t - Evapotranspiration from the water table is negligible since the water table is normally 3 m below the ground surface.

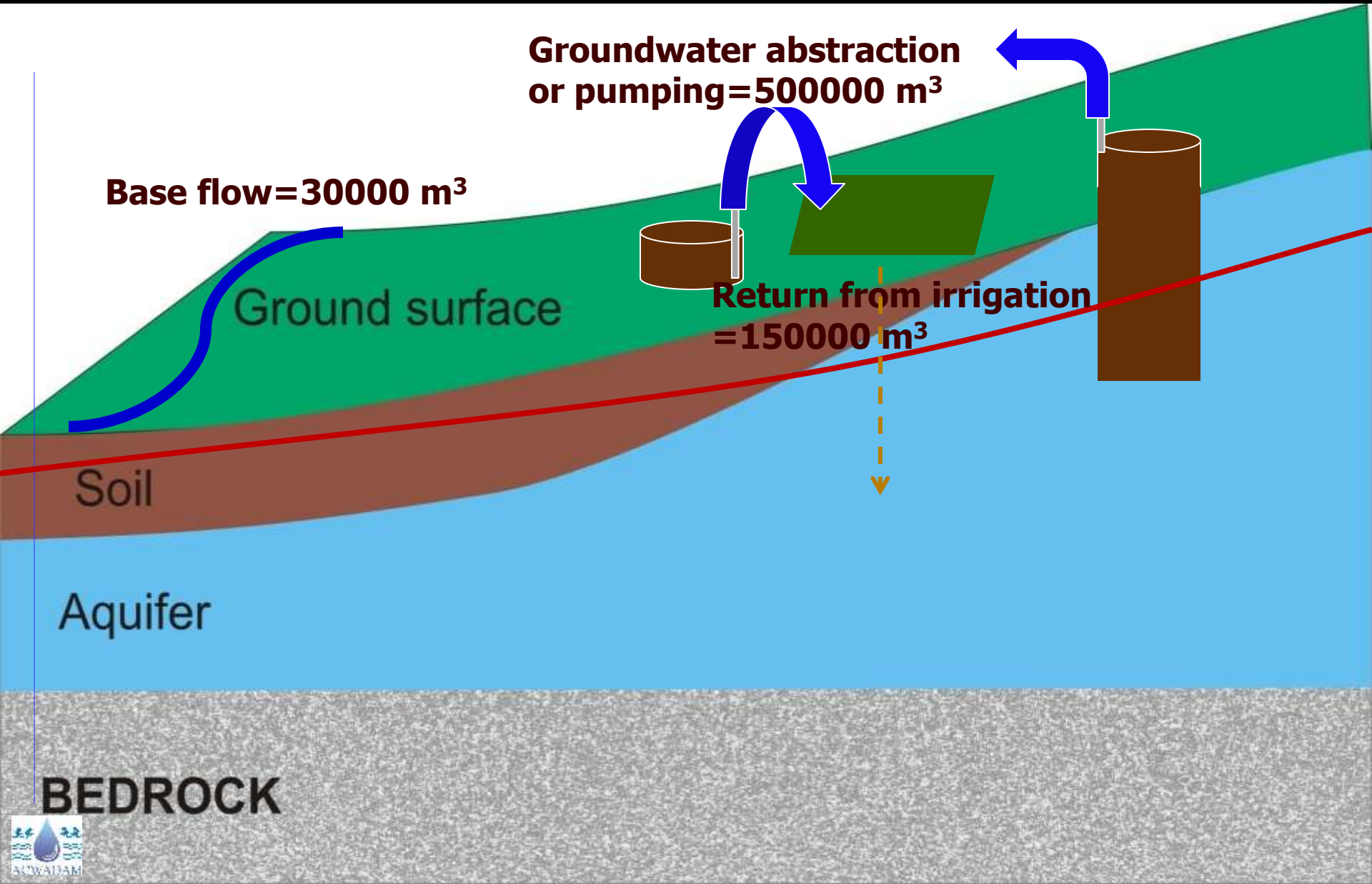


It is easier to measure “groundwater discharge”...



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Major components of groundwater discharge and return flow



A simplified groundwater balance

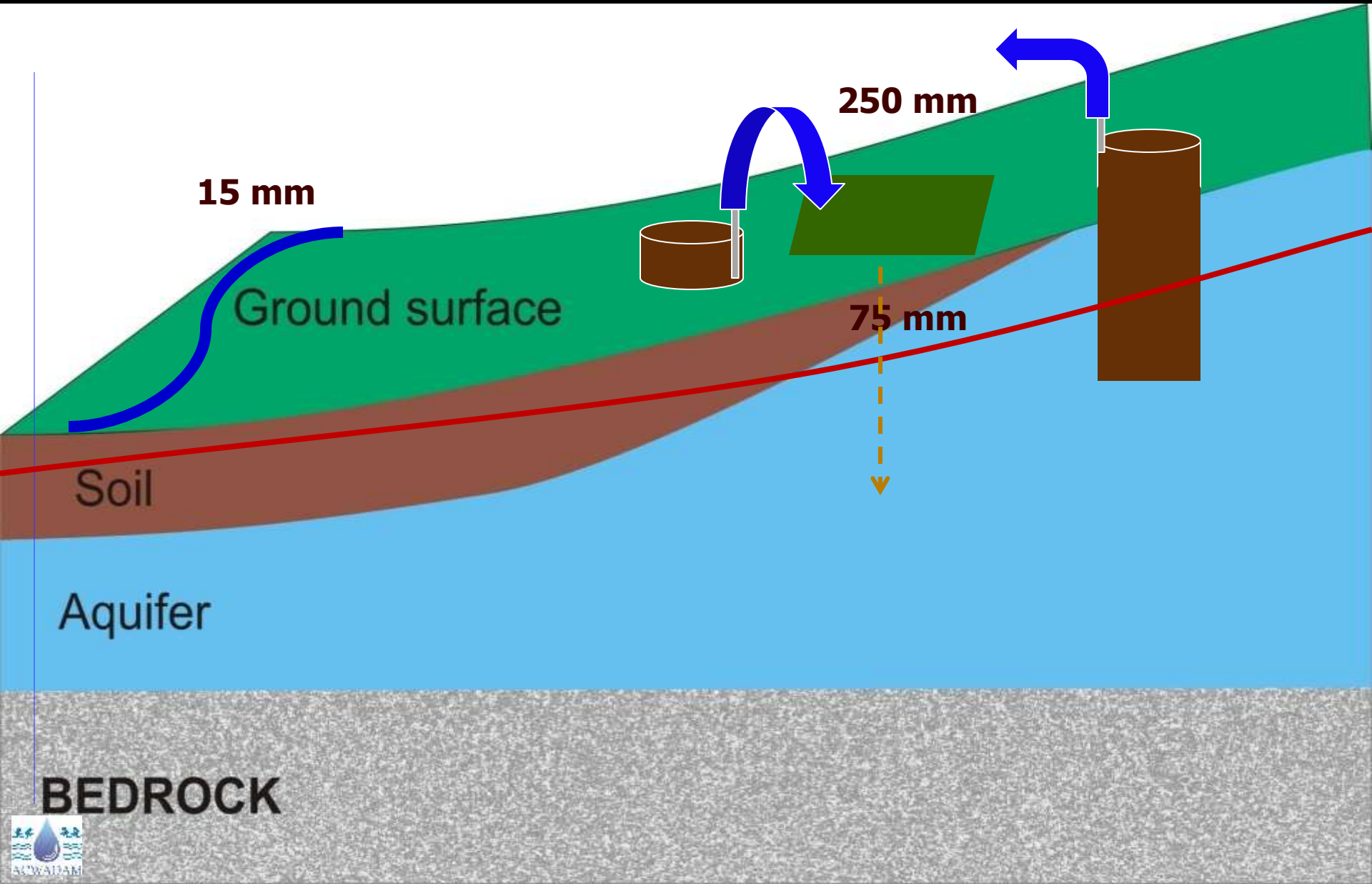
$$\Delta S = Q_p + B_f$$

- Q_p can be estimated directly using measurement of discharge from individual wells multiplied by the number of wells (in this example it is $5 \times 10^5 \text{ m}^3$ per year).
- For an area of about 200 hectares this is equivalent to 250 mm of water.
- Some irrigation water returns to wells as *Irrigation Return Flow*. In this area 30% of water applied as flood irrigation is estimated to return to the aquifer as "irrigation return flow".
- Hence, the *net* $Q_p = 175 \text{ mm}$
- Other losses such as B_f was measured and found to be some 30000 m^3 equivalent to 15 mm of water.

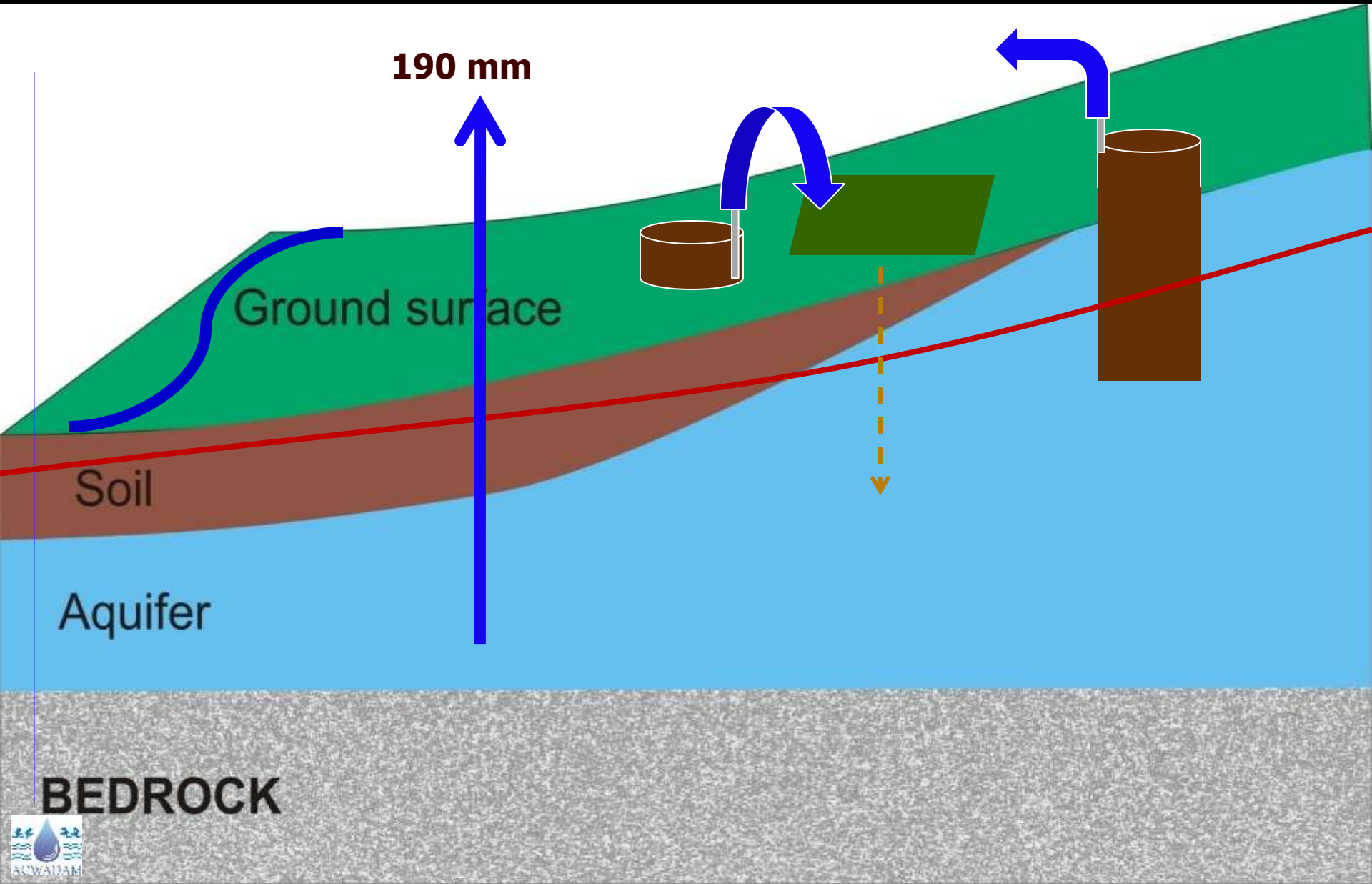
The net total discharge from the aquifer (depletion in storage) = 190 mm



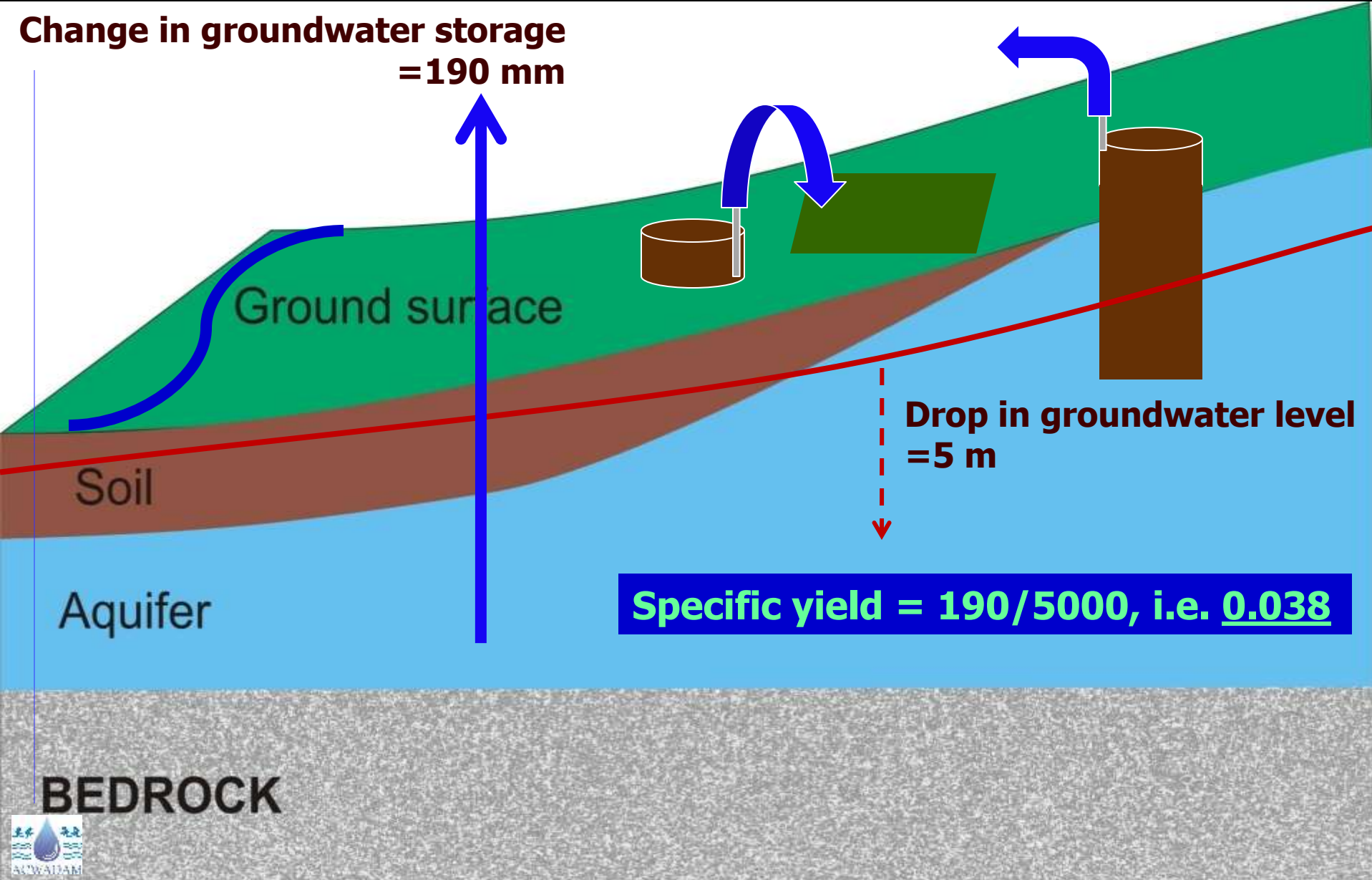
...conversion to mm for aquifer surface area of 200 mm



The processes in understanding the gw balance

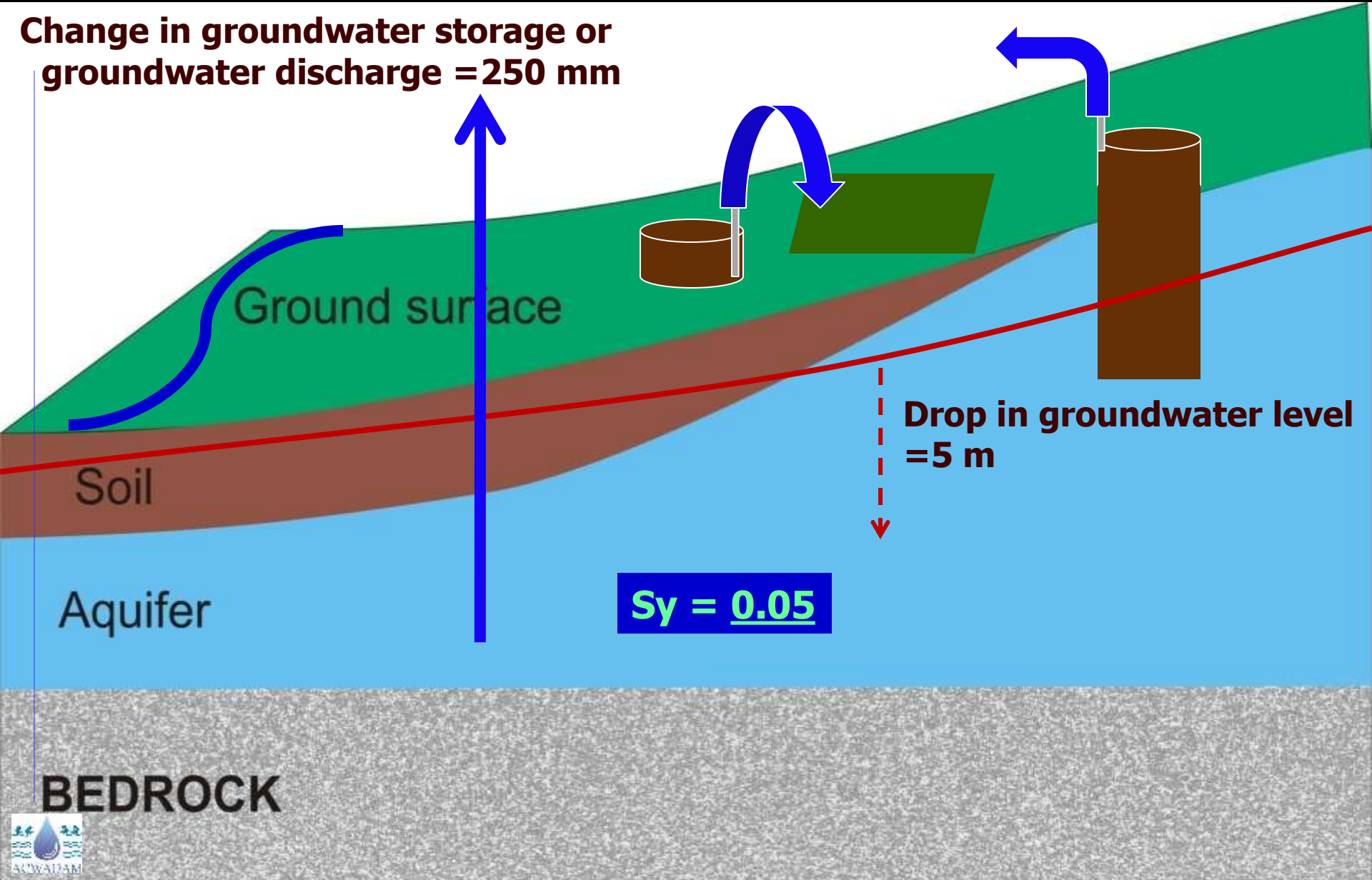


Estimating an important parameter....



Estimating another important aspect...

Change in groundwater storage or groundwater discharge = 250 mm



Using the specific yield to develop the groundwater balance further...

GW discharge = annual wl fluctuation x specific yield of the aquifer
(the wl fluctuation = 5 m)

Two things are possible now:

1. The specific yield can be estimated using:

$$Sy = \Delta s / \Delta wl$$

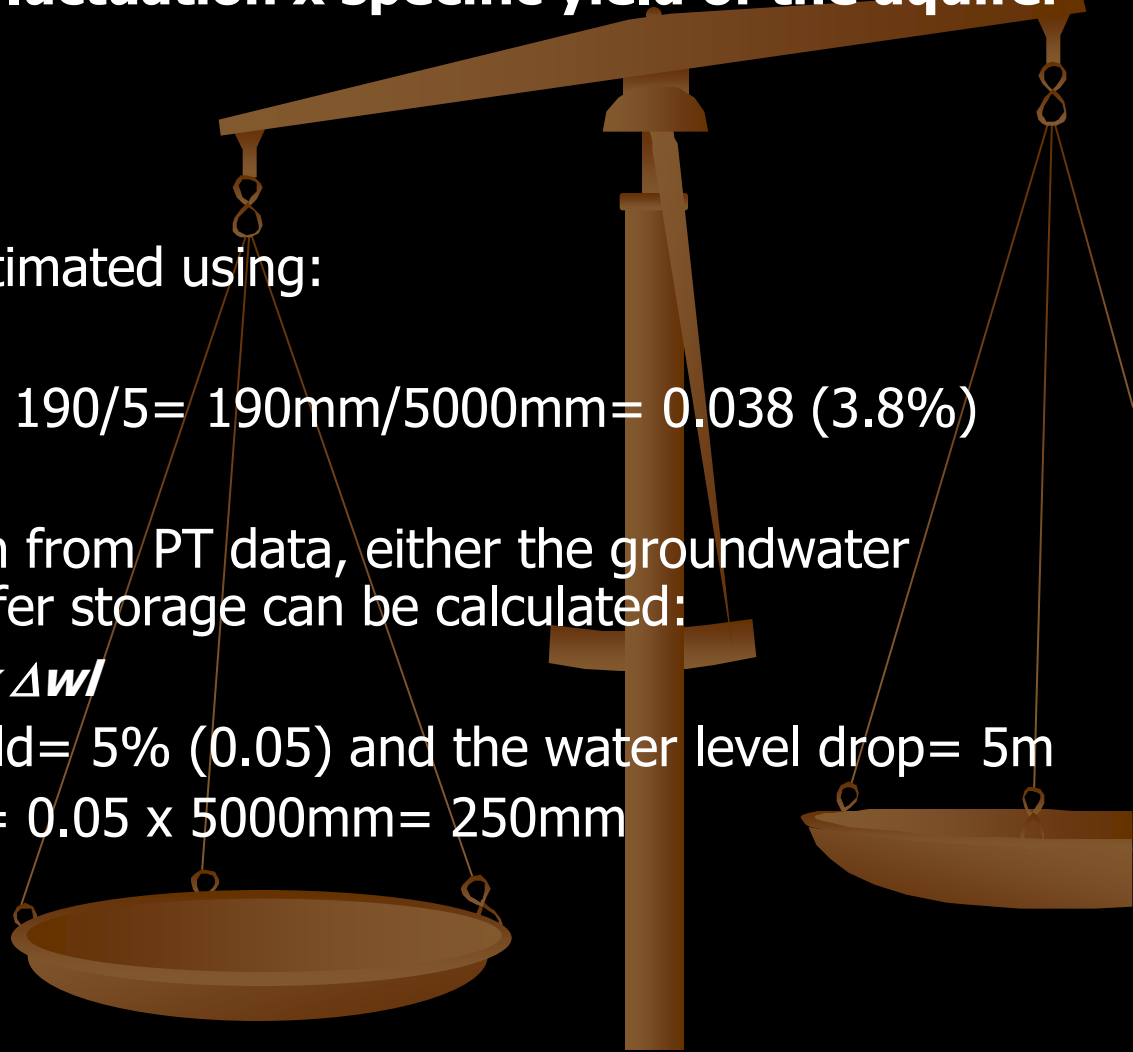
In this case the specific yield = $190/5 = 190\text{mm}/5000\text{mm} = 0.038$ (3.8%)

2. If the specific yield is known from PT data, either the groundwater discharge or loss from aquifer storage can be calculated:

$$Sy = -\Delta s / \Delta wl, \text{ i.e. } -\Delta s = Sy \times \Delta wl$$

For instance, if the specific yield = 5% (0.05) and the water level drop = 5m

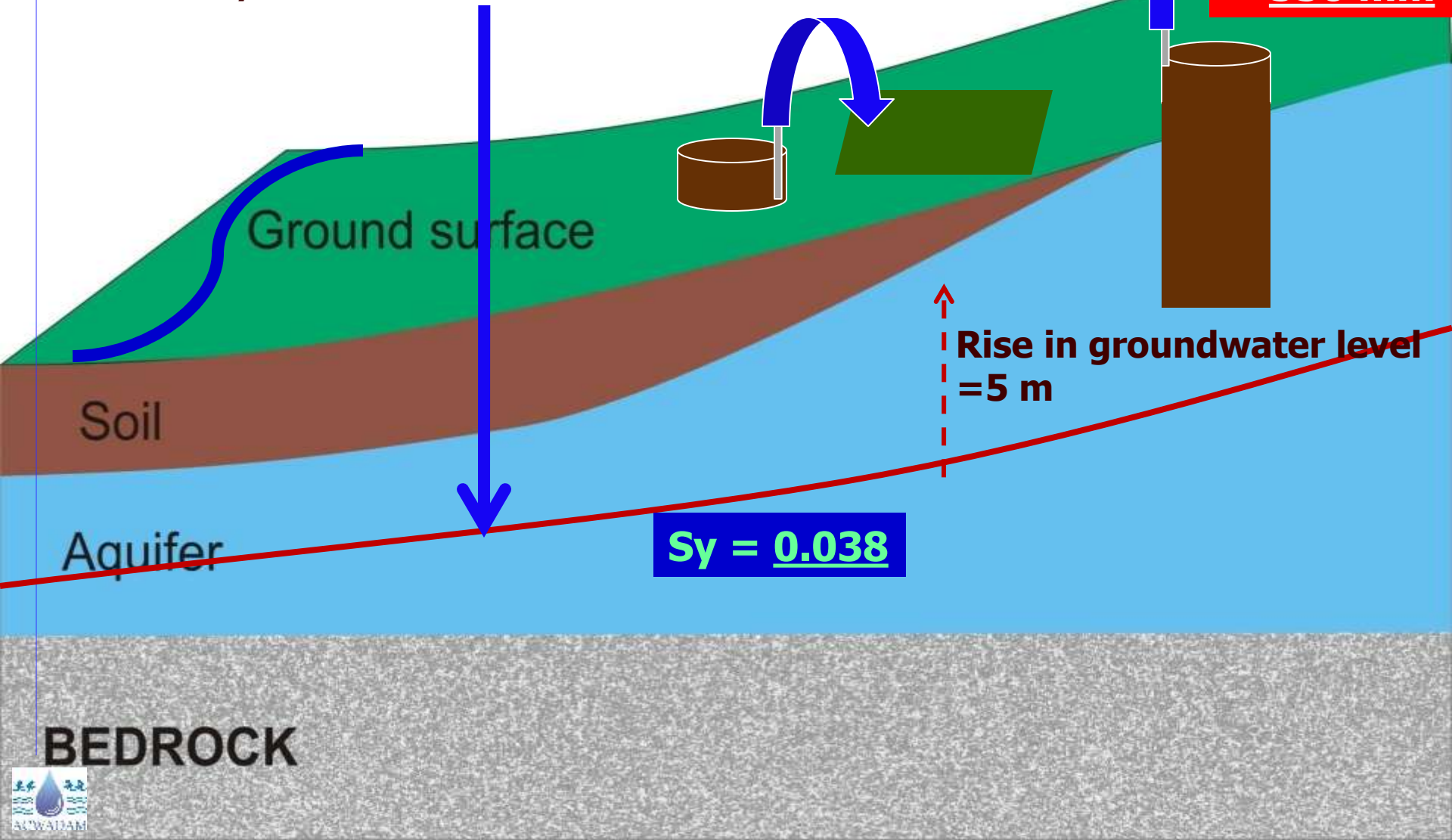
Then, groundwater discharge = $0.05 \times 5000\text{mm} = 250\text{mm}$



Recharge in proportion to rainfall...

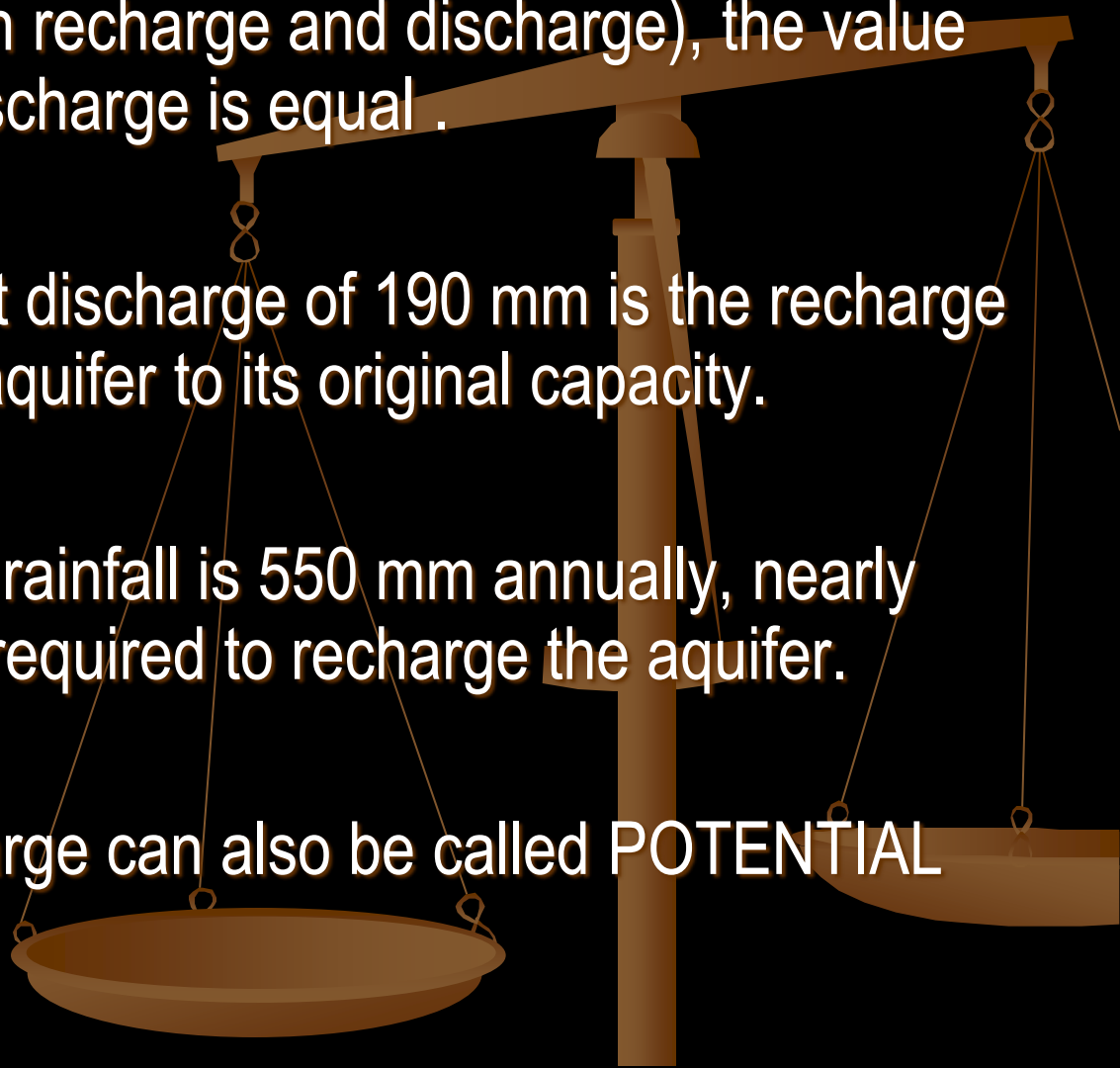
Groundwater recharge = 190 mm,
i.e. $190/550$ or 38% of rainfall

Rainfall
= 550 mm



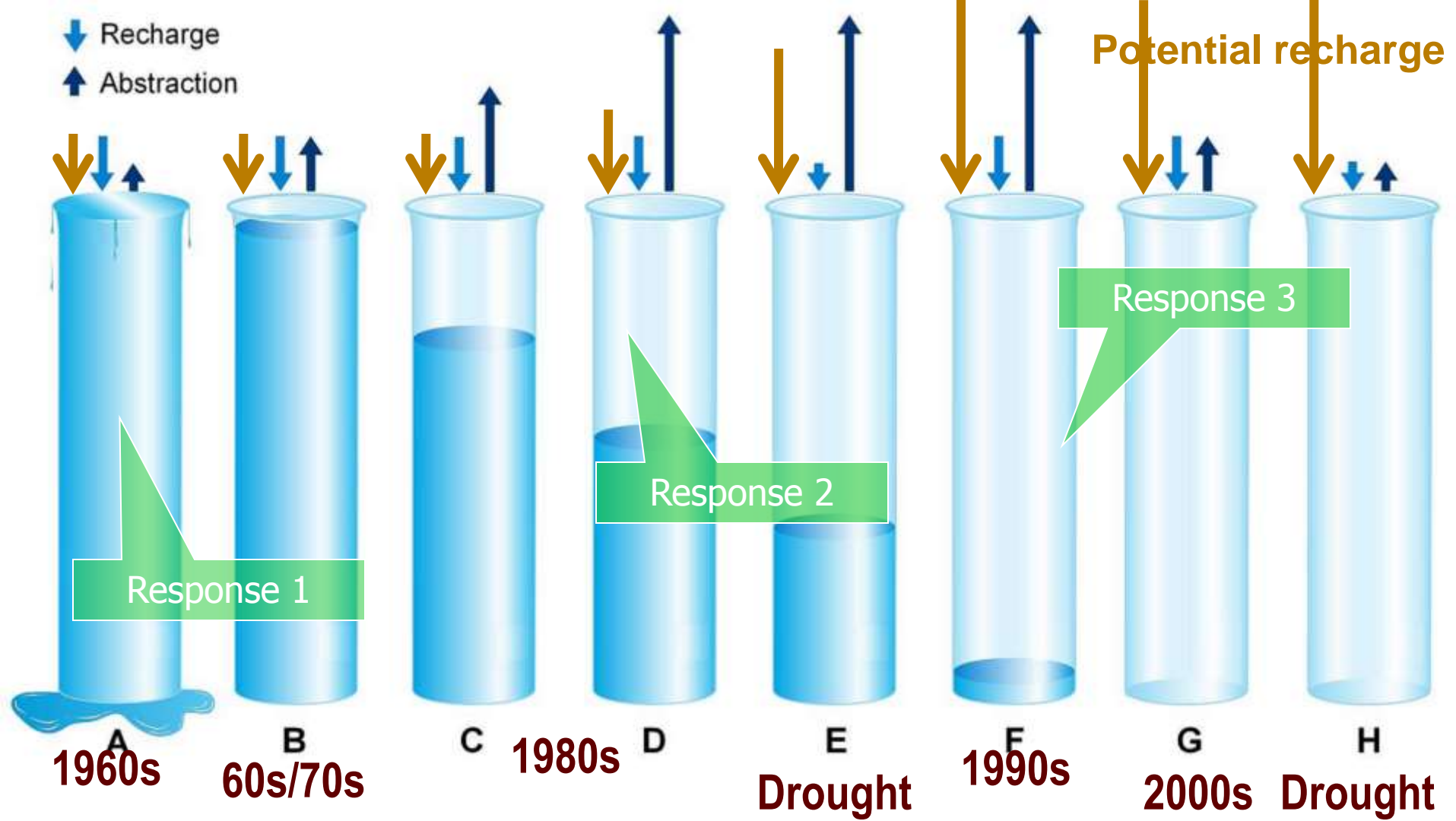
How do you equate recharge and discharge

- In areas where the aquifer empties and fills up each year (a good balance between recharge and discharge), the value of net groundwater discharge is equal .
- In other words, the net discharge of 190 mm is the recharge required to fill up the aquifer to its original capacity.
- This means that if the rainfall is 550 mm annually, nearly 34%, of the rainfall is required to recharge the aquifer.
- This estimate of recharge can also be called **POTENTIAL RECHARGE**.

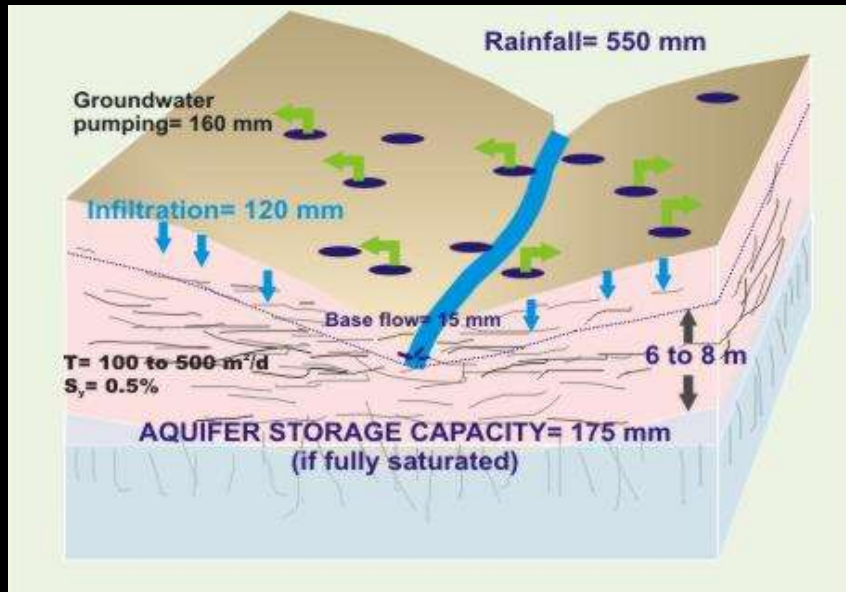


↓ Recharge
↑ Abstraction

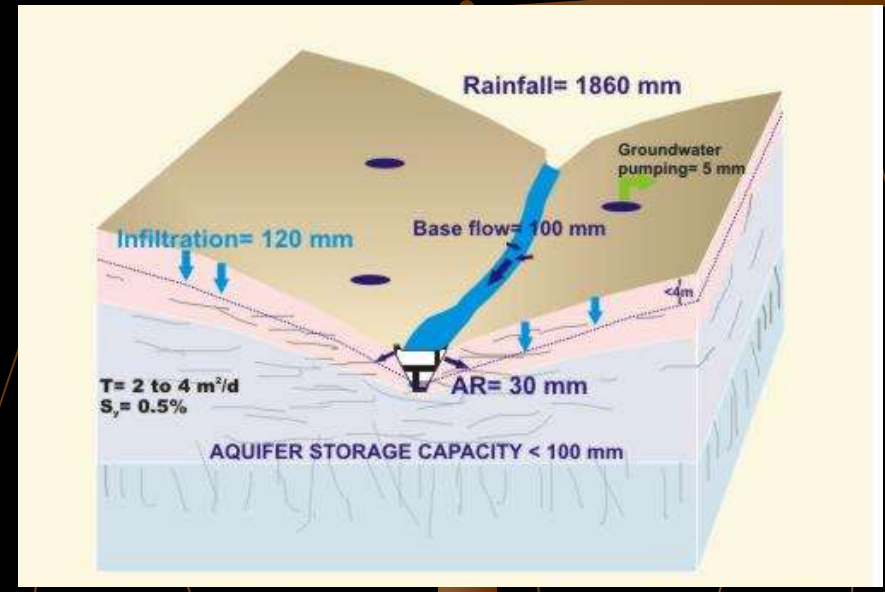
Potential recharge



Recharge is addition to the existing aquifer storage and discharge is removal from aquifer storage



Depleted aquifer storage at the end of summer



Very little depletion in aquifer storage at the end of summer

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